# **Orbital order in layered manganites probed** with <sup>57</sup>Fe Mössbauer spectroscopy

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**Abstract** In mixed-valence manganites, the doped iron species play the roles of simultaneous probes and ruiners for the orbital order occurring in the Mn-O electronic subsystem. We investigated the Mn<sup>3+</sup>-based undoped system RBaMn<sub>2</sub>O<sub>5.5</sub> and the half-doped systems RBaMn<sub>2</sub>O<sub>6</sub>, and RBaMn<sub>2</sub>O<sub>5</sub> substituting a tiny portion of Mn<sup>3+</sup> sites with Fe<sup>3+</sup>. Single-site Mössbauer spectra were observed at room temperature in the charge-orbitally ordered state, i.e., below T<sub>COO</sub>, for R=Y, Sm, Gd. Size of the rare earth ion strongly influences the quadrupole splitting in RBaMn<sub>2</sub>O<sub>6</sub>, but not in RBaMn<sub>2</sub>O<sub>5</sub>.

**Keywords** Mixed valence · Fe-doped manganites · A-site ordered perovskites · Mössbauer spectra · Effect of rare-earth ionic size on quadrupole splitting

### **1** Introduction

Charge and orbital order (COO) in manganites can melt under magnetic field, so that the material becomes metallic. Under application of very high magnetic fields the drop in resistance is colossal, however, smaller magnetic fields are sufficient to melt the less stable COO in manganites doped with Fe. In high- $T_c$  (high- $T_{COO}$ )

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manganites, Fe doping at the level of 2% suppresses the value of  $T_{COO}$  by 40 or 50 K [1]. The value of  $T_{COO}$  monotonically decreases with increasing atomic size of rare earth, but remains above room temperature for such low doping levels as 2% of Fe (cf.  $T_{COO} = 450$  K and  $T_{COO} = 350$  K in YBaMn<sub>1.96</sub>Fe<sub>0.04</sub>O<sub>6</sub> and in SmBaMn<sub>1.96</sub>Fe<sub>0.04</sub>O<sub>6</sub>, respectively). Thin films of such materials as SmBaMn<sub>2</sub>O<sub>6</sub> are expected to show the gigantic and ultrafast responses arising from COO melting that makes these materials very attractive for correlated electron device technologies [2].

In the structure of the layered mixed-valence manganites  $RBaMn_2O_{5+y}$  the largesize cations  $R^{3+}$  and  $Ba^{2+}$  form the separate layers alternating along c-axis with  $MnO_{1+y}$  layers. For the oxygen content (5+y) values of 5, 5.5 or 6 all the  $MnO_{1+y}$ layers are identical to each other. Each layer accommodates the charge and orbital order between  $Mn^{2+}$  and  $Mn^{3+}$  (y = 0),  $Mn^{3+}$  and  $Mn^{3+}$  (y = 0.5) and  $Mn^{3+}$  and  $Mn^{4+}$  (y = 1) shown in Fig. 1.

### 2 Experimental

Non-conventional multistep gas-heat-treatment procedures were used to synthesize these <sup>57</sup>Fe-doped powdered samples for R=La, Pr, Nd, Sm, Gd, Y. Integer-valued and half-integer-valued oxygen contents (y = 0, y = 0.5 and y = 1) in RBaMn<sub>2</sub>O<sub>5+y</sub> are well-defined and isolated. They are stabilized owing to the orbital ordering of three different types shown in Fig. 1. However, easy sample annealing in reducing or oxidizing gas atmospheres may not lead to these orbitally ordered states. A-site disordered perovskites would result with no superstructure. Instead, in the A-site of the parent perovskite structure, the large-size cation R<sup>3+</sup> and Ba<sup>2+</sup> layered arrangement must be formed beforehand. The layered R-Ba superstructure becomes a crucial prerequisite for the orbital order of the Mn-O network. The layer-growing



treatment in the clean reducing 6 N Ar atmosphere at 1,350°C was always necessary to obtain the A-site ordered deoxygenated structures (y = 0) at the first synthesizing step [3–5]. Subsequently, oxygenating and reducing treatments in layer-preserving regimes at low temperature ( $\sim$ 500°C) allowed varying only the oxygen contents and related orbital arrangements within Mn-O layers.

#### **3 Results and Discussion**

When a small amount of manganese is substituted with iron we observe that all Fe ions take the form of  $Fe^{3+}$ . All the  $Fe^{3+}$  ions reside in the sites of  $Mn^{3+}$ . This is not surprising because the  $Mn^{3+}$  and  $Fe^{3+}$  ions are very similar in size. In contrast, the site of  $Mn^{4+}$  is too small for  $Fe^{3+}$  and the site of  $Mn^{2+}$  is too large for  $Fe^{3+}$ .

Against the similarity of ionic sizes and charges, the orbital properties of the  $Mn^{3+}$ and  $Fe^{3+}$  ions are very different. The Jahn-Teller (JT) ion  $Mn^{3+}$  tends to sit in a low-symmetry environment. The symmetry degrades when only one of  $e_g$  orbitals is half-filled. However, for  $Fe^{3+}$  ion both of the  $e_g$  orbitals are half-filled, so that the semi-closed d-shell of  $Fe^{3+}$  is spherically symmetric. Unlike to the JT ions, such as  $Mn^{3+}$  or  $Cu^{2+}$ , the ion  $Fe^{3+}$  have no valence contribution to the electric field gradient (EFG) at the place of <sup>57</sup>Fe nuclei. Therefore, when  $Fe^{3+}$  is placed at the octahedral site of  $Mn^{3+}$  the observed quadrupole splitting reproduces faithfully the distortion of octahedral environment of  $Mn^{3+}$ .

In agreement with the number of the non-equivalent structural positions for  $Mn^{3+}$  the Mössbauer spectra are single-site in oxygen-depleted (O<sub>5</sub>) and oxygen-saturated (O<sub>6</sub>) samples, but show two doublets for the intermediate oxygen content (Fig. 2). In the latter case, the major doublet comes from the site with the in-plane orientation of the principal component of EFG. The minor subspectrum for y = 0.5, as well as the spectra for y = 0 and y = 1 come from sites with out-of-plane orientation of the main component of EFG ( $V_{ZZ} \perp MnO_{1+y}$  planes).



The 5-fold and 6-fold coordinations adopted by the  $Fe^{3+}$  ions are pyramidal in (a) and (b, minor doublet) and octahedral in (b, major doublet) and (c). Basically, they are the same as for  $Mn^{3+}$ . However, due to the difference of the orbital configurations of the  $Fe^{3+}$  and  $Mn^{3+}$  ions there appears some difference between the FeO<sub>5</sub> and MnO<sub>5</sub> pyramids [5]. Exact locations of the dopant  $Fe^{3+}$  and host  $Mn^{3+}$  inside of the pyramid do not coincide. The isotropic dopant ion  $Fe^{3+}$  is shifted towards the apex to equalize the fifth out-of-plane Fe-O distance with the four in-plane distances. In absence of such a shift we must expect much larger quadrupole splitting than observed.

Filling the in-plane orbital of the  $e_g$ -doublet in RBaMn<sub>2</sub>O<sub>6</sub> (Fig. 1, c) induces in MnO<sub>6</sub> octahedra the strong vertical contraction that is reflected in the value of quadrupole splitting. The contraction is enhanced with smaller size of the rare earth R. Therefore, the absolute value of quadrupole splitting  $\Delta E_Q$  increases when the volume of the reduced cell (perovskite-like cell) decreases (left-hand linear dependence in Fig. 3). On the other hand, in RBaMn<sub>2</sub>O<sub>5</sub>, where the thickness of the MnO<sub>2</sub> layer is determined mainly by the spherically symmetric ion Mn<sup>2+</sup>, the value of  $\Delta E_Q$  is roughly independent of the size of rare-earth ion R<sup>3+</sup> (right-hand line in Fig. 3).

Asymmetry between the doublet line intensities arises mainly from platy habitus of layered crystals whose main surfaces are perpenducular to  $V_{ZZ}$ . The out-ofplane and in-plane 3d-orbitals are the half-filled orbitals of  $Mn^{3+}$  in the  $O_5$  and  $O_6$  cases, where  $V_{ZZ} > 0$  and  $V_{ZZ} < 0$ , respectively. The opposite signs of  $V_{ZZ}$  results in opposite doublet asymmetries in Fig. 2 (a) and (c) owing to platy habits of microcrystals aligned in the plane of Mössbauer absorber. Additionally, a smaller asymmetry contribution exists owing to vibrational anisotropy. The total asymmetry is somewhat enhanced by this effect (Goldanskii–Karyagin) in  $O_6$ , but weakened in  $O_5$  case [1, 6].

The orientation of  $V_{ZZ}$  along the vertical axis of the tetragonal P4/mmm modification [5] is obviously dictated for all R (except Y) in RBaMn<sub>2</sub>O<sub>6</sub> by the axial site symmetry. The signs of  $V_{ZZ}$  are always confirmed by the ionic point charge model. Applying the ionic model to FeO<sub>5</sub> pyramid (y = 0), one expects the positive sign of  $V_{ZZ}$  as observed. For the octahedron (y = 1), strongly compressed along

vertical direction, the ionic model gives  $V_{ZZ} < 0$ . Although YBaMn<sub>2</sub>O<sub>6</sub> undergoes monoclinic distortion at T<sub>COO</sub>, this relatively small distortion has no effect on quadrupole splitting. As shown by linear fit in Fig. 3 and by our structure refinements, the vertical compression of octahedron increases quickly with decreasing the volume of perovkite cell; this strong compression continues to determine the value of quadrupole splitting even in presence of additional monoclinic distortion.

The fact that only one  $Mn^{3+}$  site exists in the cases of  $O_5$  and  $O_6$  Mössbauer spectra shows that the charge-ordered domains are configured under control of the frozen disorder of Fe. Random distribution of the Fe dopants tends to break down the ordered domains; however, the charge order would be reconstructed, centering all  $Mn^{3+}$  sites at Fe<sup>3+</sup> ions. In contrast, two distinct  $Mn^{3+}$  sites exist in the  $O_{5.5}$  phases synthesized at 1,450°C [5]. On cooling the iron dopants fall out preferentially into the octahedral sites of GdBa $Mn_2O_{5.5}$ , so that the population of pyramidal Fe species (inner doublet area) is found to be 4 times smaller.

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